PALEONTOLOGICAL
RESEARCH INSTITUTION

The
Teacher-Friendly
Guide™
to Climate Change

CLIMATE CHANGE TOOLKIT
The Carbon Cycle: Respiration

Equipment:

- Small flowerpots filled with biologically active soil (not sterile potting soil); 1 or 2 per group.
- Optional: a way to chill or warm the soil pots; bags of ice, a refrigerator, a hot plate or warming mat, etc.
- One-liter soil flux chambers (1 or 2 per group)
- Duct tape (or similar, to make an opaque chamber)
- Lab CO_2 probes and necessary hardware/software for probes
- Soil thermometers (optional)

Introduction

Measure the flux of carbon dioxide from soil to the atmosphere using a lab CO_2 probe and home-made flux chambers, then graph the results.

Why?

Photosynthesis transfers carbon dioxide from the atmosphere to plants, releasing oxygen in the process. Respiration runs this reaction backwards, releasing CO_2 while consuming oxygen.

$$C_6H_{12}O_6 + 6O_2 = 6CO_2 + 6H_2O$$

glucose + oxygen = carbon dioxide + water

Globally, photosynthesis removes 120 Pg (120 billion metric tons) of carbon every year from Earth's atmosphere and turns it into the carbohydrates that make up living plants. Respiration and decay within ecosystems return that *same amount* of carbon to the atmosphere every year. We can quantitatively measure this flux in any outdoor environment, and compare it with anthropogenic carbon fluxes. Global climate change - especially rising temperature - can perturb the carbon cycle in both the short and long term.

What?

The movement of carbon dioxide from the atmosphere into biomass is part of the global carbon cycle (Figure 1).

- On very long timescales (thousands to millions of years) the geologic carbon cycle moves carbon from the Earth's mantle and crustal rocks into the atmosphere through volcanic eruptions. Carbon is removed from the atmosphere by chemical reactions that weather and decompose rocks.

- On shorter timescales carbon moves from biomass to the atmosphere and back via photosynthesis and respiration.
- In the ocean carbon also cycles to the atmosphere and back through exchange with sea water.

Each of these three parts of the carbon cycle runs in a steady state, with all the inputs and outputs in balance. The fluxes can be very large, but as long as they are all balanced there is no overall accumulation in any one place. In contrast, the human contribution to the carbon cycle is one-way: humans burn fossil fuels and biomass releasing CO_2 to the atmosphere. There is no equivalent return process, thus humans upset the natural balance.

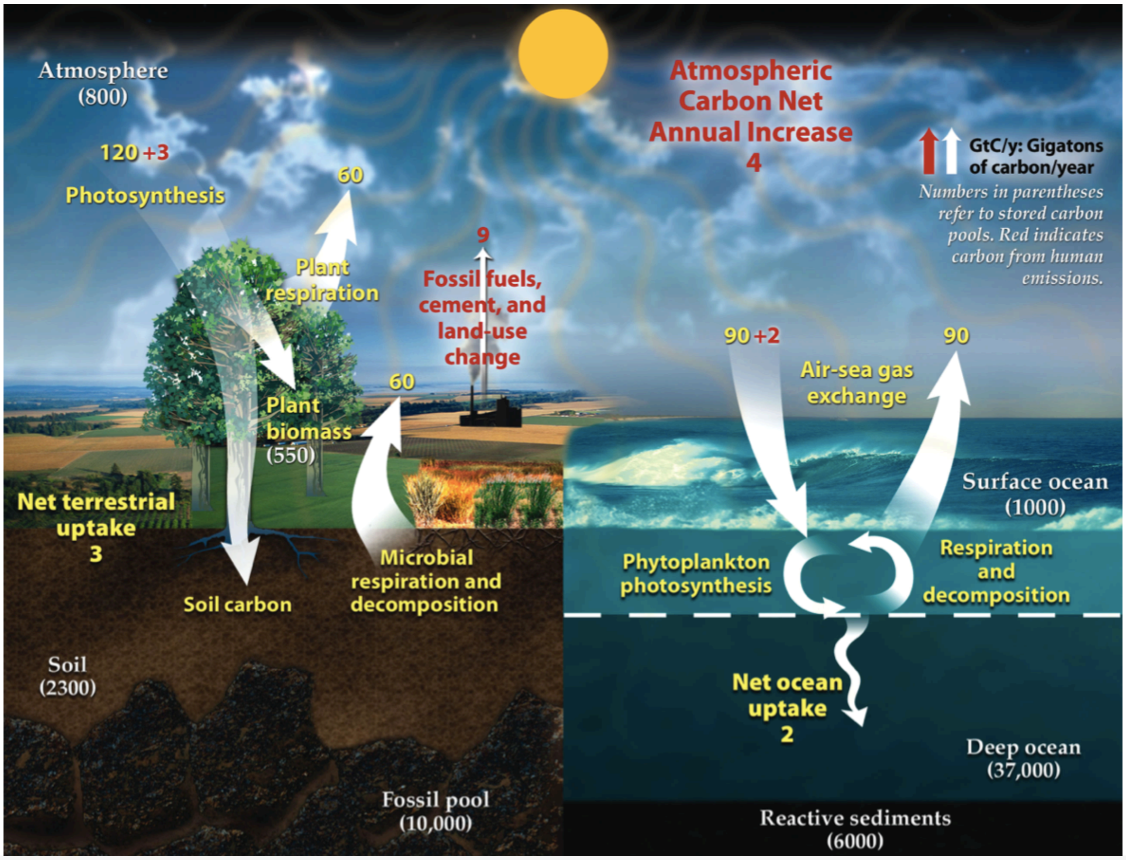


Figure 1: This diagram of the fast (non-geologic) carbon cycle shows the movement of carbon between land, atmosphere, and oceans. Yellow numbers are natural fluxes, and red are human contributions in gigatons of carbon per year. White numbers indicate stored carbon. https://earthobservatory.nasa.gov/Features/CarbonCycle/images/carbon_cycle.jpg

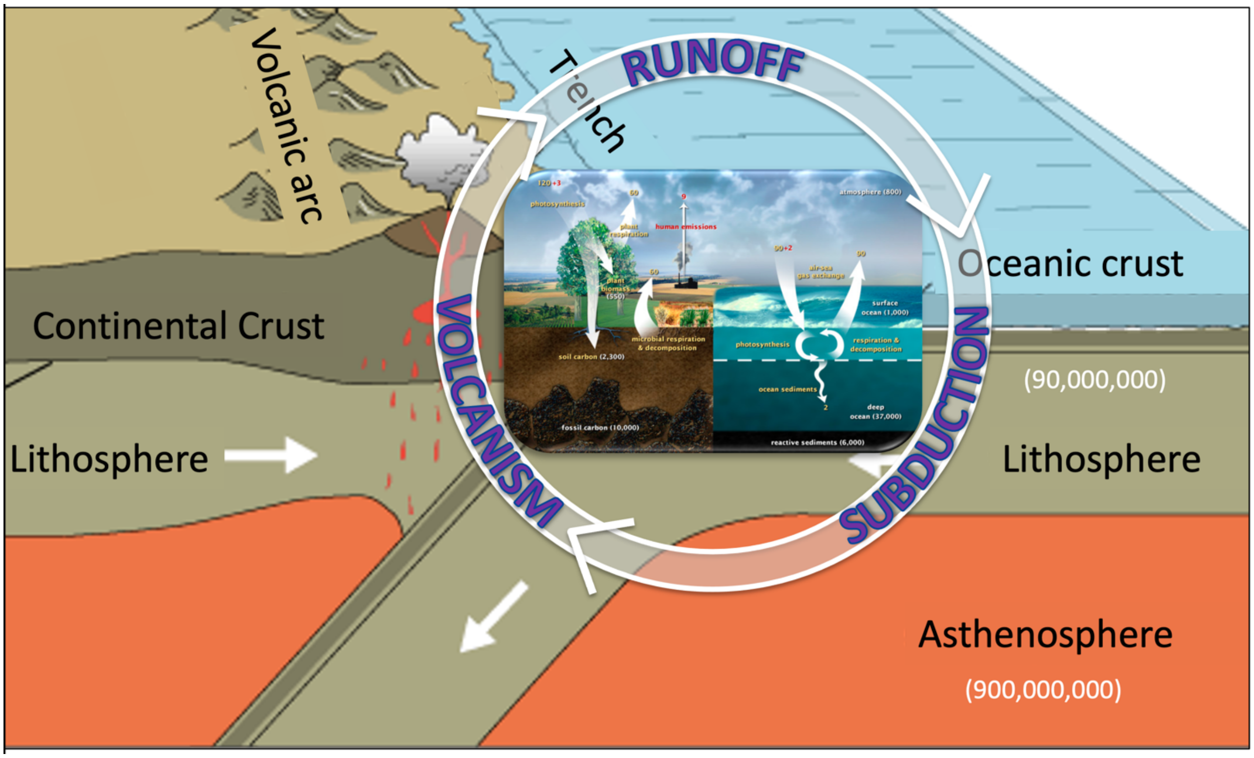


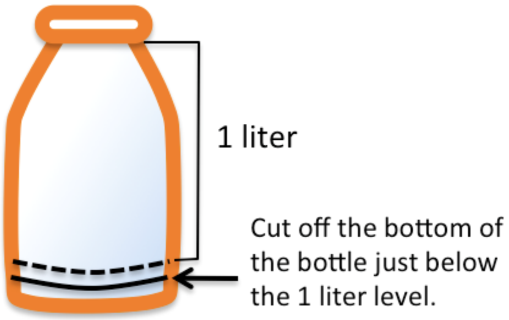
Figure 2: The geologic carbon cycle. The terrestrial and marine carbon cycles from Figure 1, above, are placed within the plate tectonic setting of the geologic carbon cycle. Geologic processes are slower, but move many more gigatons tons of carbon (units are the same as above).

How?

We will measure the respiration and decomposition flux from the soil using a CO_2 probe. From these data we can estimate the contribution of our local environment to the global carbon cycle.

PART 1 - BUILD THE EXPERIMENTAL COMPONENTS

Make a plastic juice bottle flux chamber. The flux chamber will control the movement of CO_2 gas between the soil and probe, and allow us to make quantitative measurements. Choose a 2-quart (or similar) plastic juice bottle with a wide enough mouth to fit the CO_2 probe. Pour in one liter of water (that you've pre-measured in a beaker or graduated cylinder). With the cap on, turn the bottle upside down and draw a line at the one liter mark. Discard the water and cut off the bottom of the juice bottle evenly, just below the one liter line.



This will allow us to push the flux chamber a little bit into the soil to give it some stability during our measurements. Measure the cross-sectional area of the opening at the bottom of the bottle (*length* x *width* for a rectangular bottle; or πr^2 for circular). Measure in *meters* and record the value (or measure in cm and convert to meters: 1 cm = 0.01 m). Finally, cover the bottle with duct tape (or paint, or any other opaque material so that light does not pass through the bottle).

PART 2 – SET UP AND RUN THE EXPERIMENT

Prep your CO_2 probe as you would for any experiment (set up computer interface, etc.). We will record CO_2 values at a time interval of 10 seconds over an experimental run of about 5-10 minutes.

➢ See instructions on p.6 if you will run extra experiments with different soil temperatures.

- Remove the cap from the juice bottle flux chamber and insert the probe into the top of the bottle. The probe should fit loosely; make sure there is an air gap that will allow the CO_2 to flux through and out of the chamber. Allow a few minutes for the probe to equilibrate.
- Press the flux chamber into the soil so that the one-liter line is at the soil level.
- Collect data for 5 – 15 minutes. Note that at the beginning of the experiment the CO_2 data may be noisy. If so, when analyzing the data ignore these first few minutes of measurements.

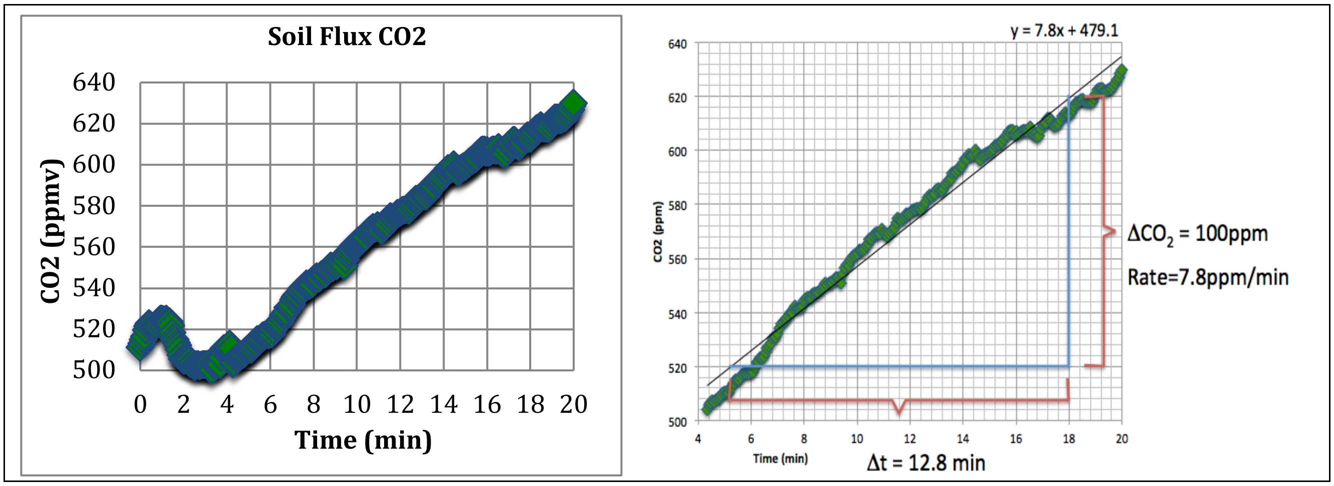


Figure 3: (Left) Sample data; a time series of carbon dioxide measurements. Note that data are noisy at the beginning of the experiment. For these data the slope is calculated only from 4-18 minutes. For any data set, look for a consistent increase early in the experiment; these will be the most characteristic data, unaffected by later CO_2 buildup in the chamber. (Right) Calculation of rate of change for CO_2 *data.*

PART 3 – ANALYZE THE DATA

The data collected by the CO_2 probe is in units of concentration – parts per million. We want to work in units of mass; grams, kilograms or metric tons of CO_2. For example, in Figure 1 the fluxes are given as gigatons of carbon per m^2 per year. In the chart below we will convert the measured concentration flux into a mass flux of CO_2 – the mass of CO_2 per unit volume per unit time. We will use metric (mks) units, so we should end up with kg/s (although in this case its OK to use minutes if the probe was set up that way).

- Graph the data: time (min or sec) on the X-axis, and CO_2 (ppmv) on the Y-axis. To calculate the flux we will find the rate of change of the CO_2 concentration in the one-liter flask measured over a fixed interval of time (Figure 2).
- Draw a line – or use a computer to fit a straight line – through your data and calculate the slope. This is the flux rate in units of ppmv/time.

UNIT CONVERSION FROM CONCENTRATION TO MASS FLUX OF CO_2

Steps in Unit Conversion	**Result**
1. Calculate the cross-sectional area of the flux chamber in units of square meters.	m^2
2. Recall that at standard temperature and pressure (STP) one mole of gas occupies a volume of 22.4 liters. Invert this quantity to find the number of moles in a single liter of gas (e.g. the size of our flux chamber).	mol
3. From the graphed data, calculate the rate of change of CO_2 in the flask in ppmv/second.	ppmv/s
4. Convert this rate of change from ppmv to its decimal equivalent (to do this, divide the rate by 1,000,000). This allows us to use the rate to calculate mass.	
5. Multiply the rate from Step 4 by the number of moles from Step 2, then divide by the cross-sectional area of the flux chamber from Step 1. The result is the flux of CO_2 in units of mol/m^2/s.	mol/m^2/s
6. The mass of CO_2 is 44 g/mol (C=12, O=16). Multiply the result from Step 5 by 44 g/mol to convert the flux to units of g/m^2/s.	gCO_2/m^2/s
7. This result is what we were looking for: the mass flux of CO_2 per unit area per unit time. In order to compare our result with published data, there are most likely more unit conversions to do – but simpler ones – converting g to kg or minutes to years. See below.	

Flux is an extensive property; one that changes with the size of the system. For example, the longer we observe, the more CO_2 moves through the flux chamber, or, the larger the chamber the more CO_2 will move through it. Therefore we need to normalize our measurement with respect to both time and the cross-sectional area of the flux chamber. The rate calculation already accounts for the time of observation. We still need to account for the basal area of the flux chamber.

OPTIONAL: THE RELATIONSHIP BETWEEN TEMPERATURE AND RESPIRATION

By conducting respiration experiments with small pots of soil we have the opportunity to chill or heat the soil and measure the respiration flux at different soil temperatures. This can be done two ways:

1. A side-by-side experiment with one heated pot and one chilled pot (ex, use a water bath to control temperature, making sure that the pots are sealed and cannot become water-logged)
2. A sequential experiment with one pot that begins chilled and is then warmed (as would happen during the course of a day, or seasonally)

The second experiment is described here. Store the soil pots in a refrigerator until you are ready to use them. Have all the experimental apparatus ready to go, as the soil samples will begin to warm as soon as they are removed from the fridge. If soil thermometers are available (or any metal-stemmed thermometer) they can be placed in the pots - outside the flux chamber - to monitor the soil temperature throughout the experiment (see photo).

- Remove pots from fridge and immediately measure the CO_2 flux, ca. 10 minutes (pots can be wrapped in a dish towel or other insulator to slow the initial rate of warming).
- After 10 minutes place the pots on a hot plate (or in a hot water bath; make sure pots are sealed so water cannot enter) and gently heat for 10 minutes, measuring the CO_2 flux while heating.
- Optional; heat for an additional 10 minutes while continuing to measure CO_2.
- Graph the data. This can be done as two separate graphs for heating and chilling, or as a single time series for the entire length of the experiment.
- Find the slope of the change in CO_2 concentration as a function of time for the chilled soil, and separately, find the slope for the warmed soil.

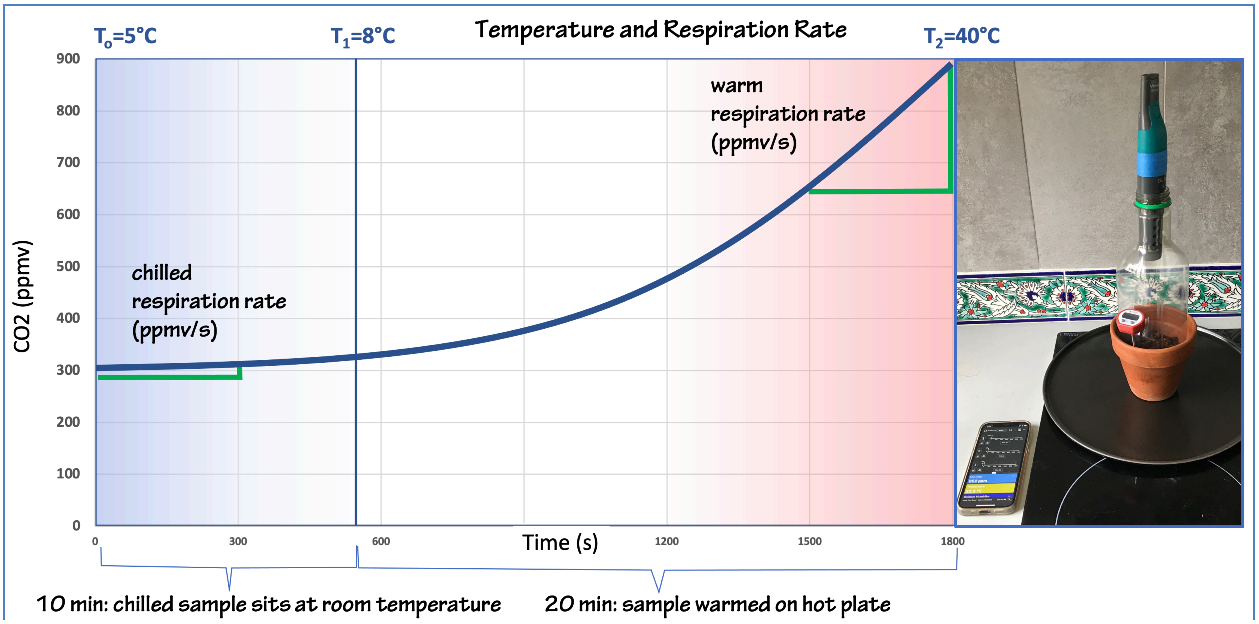


Figure 4: Example of a 30-minute experiment where a chilled soil sample is removed from the refrigerator and allowed to sit at room temp for 10 minutes while CO_2 concentration is measured. A soil thermometer in the pot records the change in temperature over this interval, from 5C - 8C. The sample is them warmed on a hot plate for 20 minutes, increasing the soil temperature to 40C. The respiration rate (slope) is calculated for the chilled soil and for the warm soil.

PART 4 – PUTTING RESULTS IN CONTEXT

Compare the data we collected with the maps of global respiration CO_2 flux (Figure 5). These data are the average annual soil fluxes in units of megagrams of carbon (not CO_2) per square kilometer per year, so for direct comparison we will need to do additional unit conversion:

8. First change grams of CO_2 to grams of carbon. The fraction of C in CO_2 is 12/44 (mass of CO_2=44, mass of C=12). Multiply your result in Step 6 by (12/44).	gC/m^2/s
9. Convert the time from seconds to years. 1yr=(365days)(24hrs)(60min)(60sec).	gC/m^2/yr
10. Convert m^2 to km^2 (1km=1000m)	gC/km^2/yr
11. Convert g to Mg (1Mg=1,000,000g)	MgC/km^2/yr

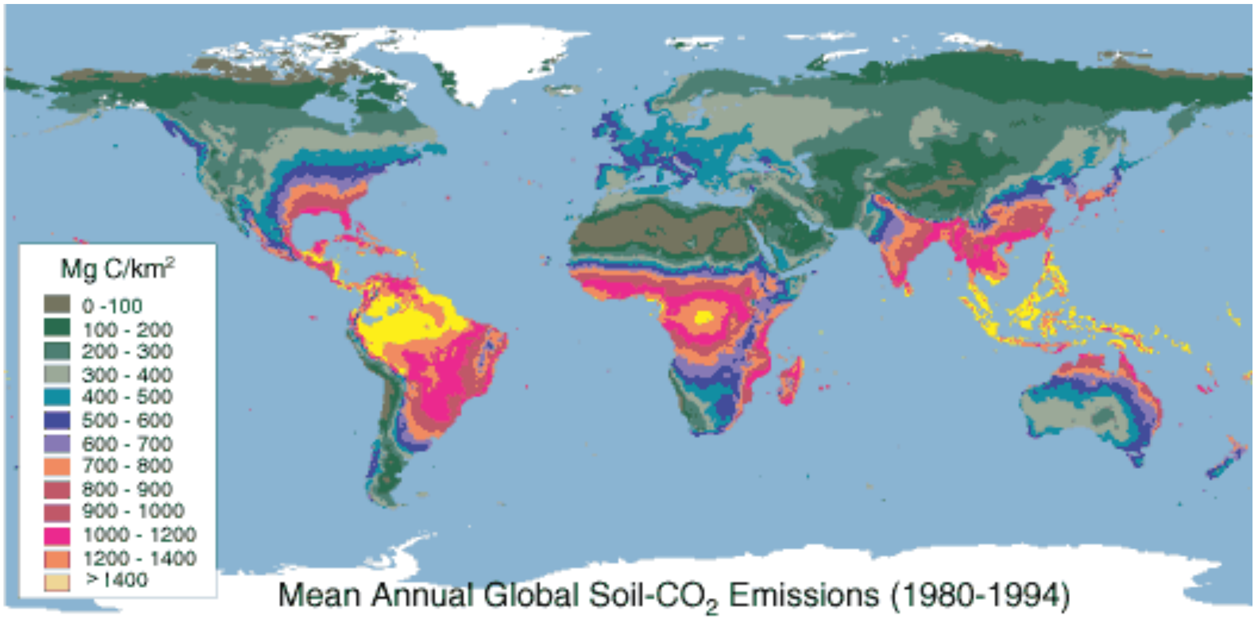


Figure 5: Global mean annual soil carbon flux (MgC/km2/yr), Raich et al., 2003 (https://cdiac.ess-dive.lbl.gov/epubs/ndp/ndp081/map_yrmn.html)

Part 5: Discussion

1. Are there any differences in the respiration rates in your experiment when the soil is warm or cold? Describe.
2. Look at the map of global soil respiration (Figure 5). Is there a relationship between your heating/chilling experiments and the pattern on the map? Describe.
3. If climate is warming globally, how might higher temperatures change this map pattern several decades into the future?
4. Look again at the global carbon cycle in Figure 1. Consider the three terrestrial reservoirs for carbon:
   - Atmosphere 800 Gt
   - Biosphere 550 Gt
   - Soil 2300 Gt

   Make a hypothesis that predicts how the rate of respiration will change in a global warming scenario, and how that will impact the size of the Atmosphere and Soil reservoirs for carbon.